

Water Movement in Soils

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- 1) Soil Forming Factors
- 2) Primary Properties that affect Water Movement in Soils
- 3) Soil Water Concepts
- 4) Where to get additional information

Today's Talking Points

- Parent Material
- Climate
- Topography
- Biological Activity
- Time

Soil Forming Factors

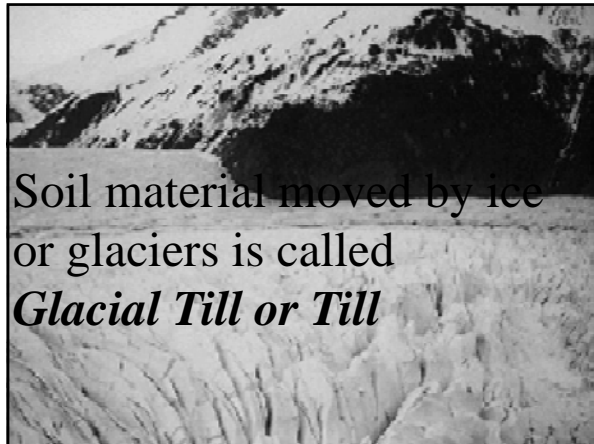
- Those that formed in place - residuum
- Those that have been transported - In Missouri: transported by Ice, Water, Wind and Gravity

Two Categories of Parent Material

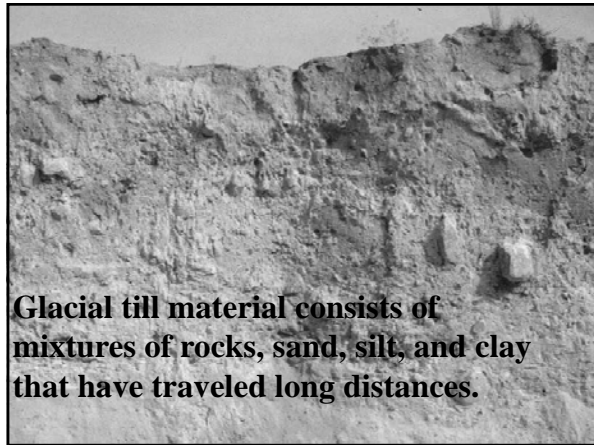
Soil that form in place from the weathering of the underlying rock or minerals is called

Residuum

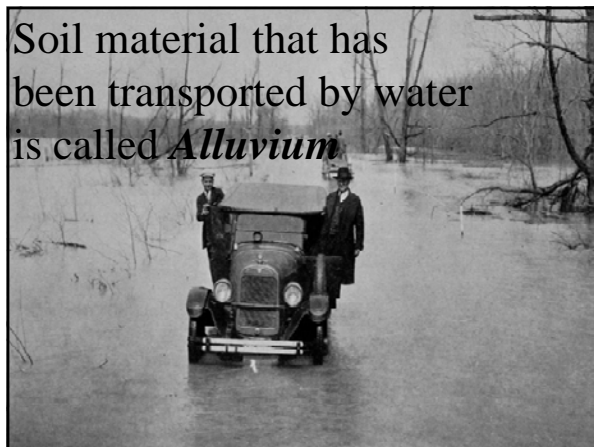




Soil material moved by ice or glaciers is called ***Glacial Till or Till***



Glacial till material consists of mixtures of rocks, sand, silt, and clay that have traveled long distances.



Soil material that has been transported by water is called ***Alluvium***

Slow velocity
flooding
produces clayey
soils.

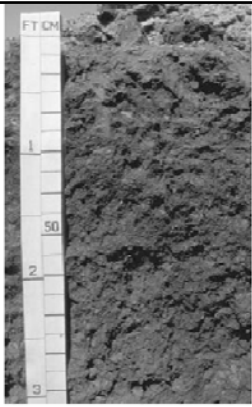
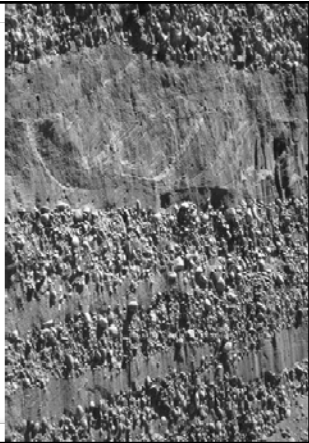


Figure 23—Typical profile of Otago silty clay loam, 6 to 1 percent slopes, occasionally flooded.

High velocity
flooding
produces
loamy soils.



Floods can move great
amounts of soil material.

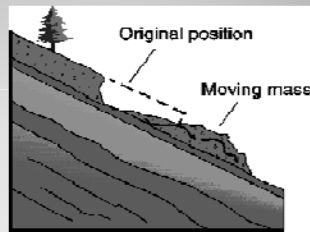


- Wind deposited - the further from the source the loess gets thinner and has more clay (less silt)
- Thickest deposits are adjacent to the Missouri and Mississippi River flood plains.

Loess



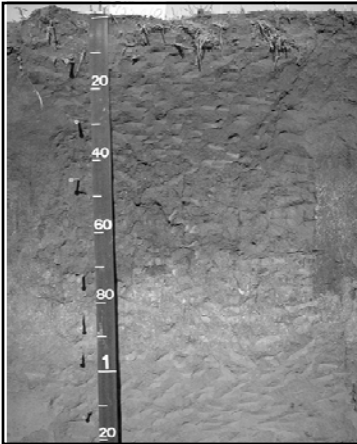
Colluvium is material transported down slope by gravity.



• Plants
• Animals

Biota

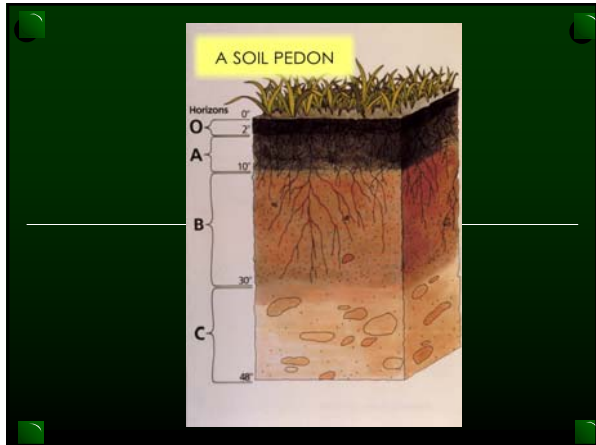




**Makes
thick dark
Surface
Horizons**

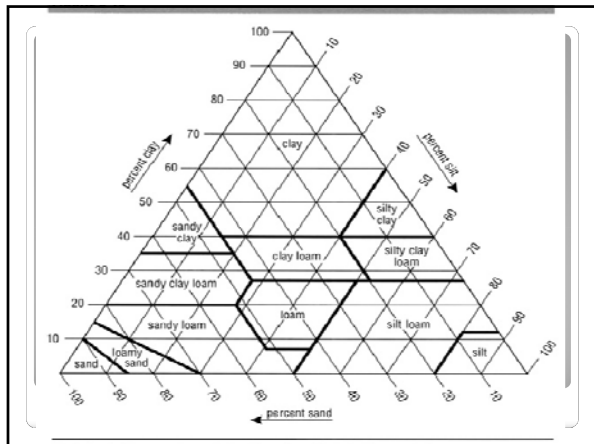
- Recent floods to "old" residuum

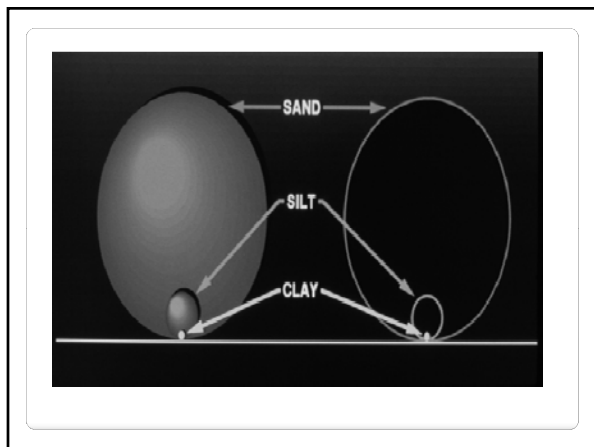
Time



- 1) Soil Texture – particle size distribution
- 2) Soil Structure
- 3) Organic Matter Content
- 4) Landform – hill slope profile

Primary Factors that influence Water Movement through the Soil





Particle Sizes

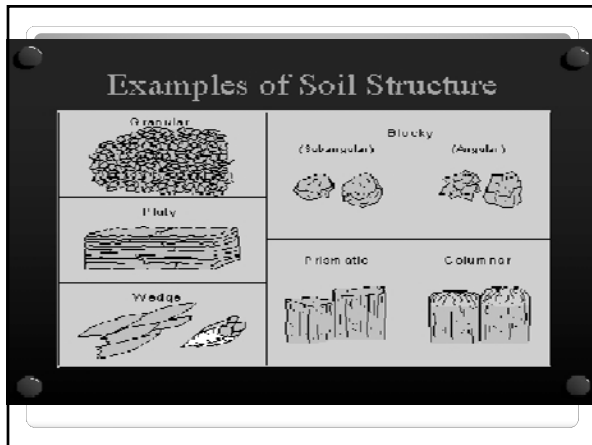
Sand: 2mm to .05mm

Silt: .05 mm to .002 mm

Clay: smaller than
.002mm

- Blocky or granular - surface layer, A horizons
- Platy or blocky - subsurface, E horizons if present
- Blocky and/or Prismatic - subsoil, B horizons (compound: prismatic parting to blocky)
- Massive or Single grained - substratum, C horizons

Structure



- Soil Organic Matter Content higher levels:
 - Will increase the soils ability to store and give up water plants.
 - Will increase water movement rates through the soil.

Soil Organic Matter

- How we can describe shapes of the land surface.
- Slope shape and gradient are important considerations.

Landforms

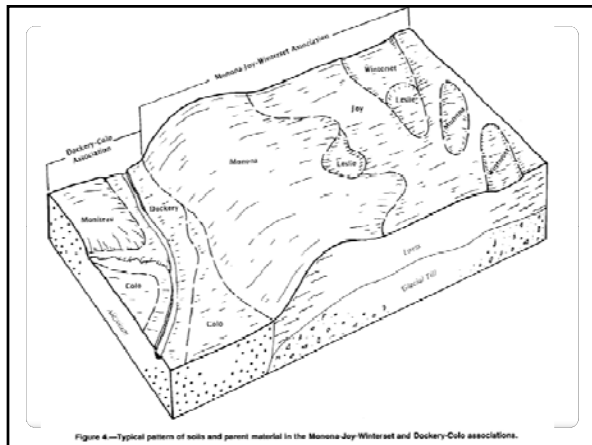


Figure 4—Typical pattern of soils and parent material in the Manana-Jay-Winter set and Dockery-Cole associations.

- In soil science, permeability is defined qualitatively as the ease with which gases, liquids, or plant roots penetrate or pass through a soil mass or layer. It is measured in length with no time component.
- Saturated hydraulic conductivity is a quantitative measure of a saturated soil's ability to transmit water when subjected to a hydraulic gradient. It can be thought of as the ease with which pores of a saturated soil permit water movement. It is measured in both length and time.
- Permeability classes and most references to permeability have been removed from the *NSSH* and replaced with the saturated hydraulic conductivity classes of the 1993 *Soil Survey Manual*.

K_{sat} Saturated Hydraulic Conductivity

SATURATED HYDRAULIC CONDUCTIVITY - PERMEABILITY			
K _{sat} Class		Permeability Class	
705.00	100.00	705.00	100.00
VERY HIGH	141.14	VERY RAPID	20.00
		RAPID	6.00
HIGH	14.17	MODERATELY RAPID	2.00
		MODERATE	0.60
MODERATELY HIGH	1.417	MODERATELY SLOW	0.20
		SLOW	0.06
MODERATELY LOW	0.1417	VERY SLOW	0.0015
LOW	0.01		
VERY LOW	0.00		
$\mu\text{m}/\text{sec}$	in/hr	$\mu\text{m}/\text{sec}$	in/hr
$\mu\text{m}/\text{sec} \times 0.1417 = \text{in}/\text{hr}$		$\text{in}/\text{hr} \times 7.0372 = \mu\text{m}/\text{sec}$	

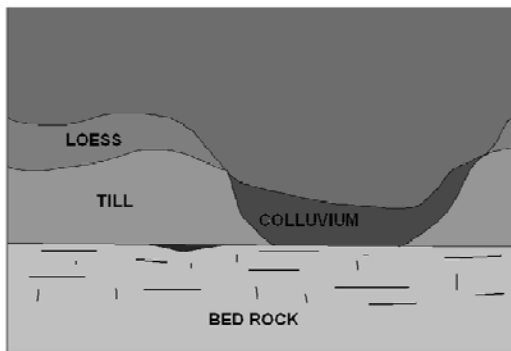
- Saturated hydraulic conductivity is the amount of water that would move vertically through a unit area of saturated soil in unit time under unit hydraulic gradient.
- If you find old permeability information, you can convert from $\mu\text{m}/\text{sec}$ to in/hr by multiplying $\mu\text{m}/\text{sec}$ by .1417

- Do you have an awareness of the factors of soil formation?
- How about the primary factors that influence water movement through soil?
- What is Saturated Hydraulic Conductivity?
- Ready to move on to Soil Water Concepts?

Quick Review

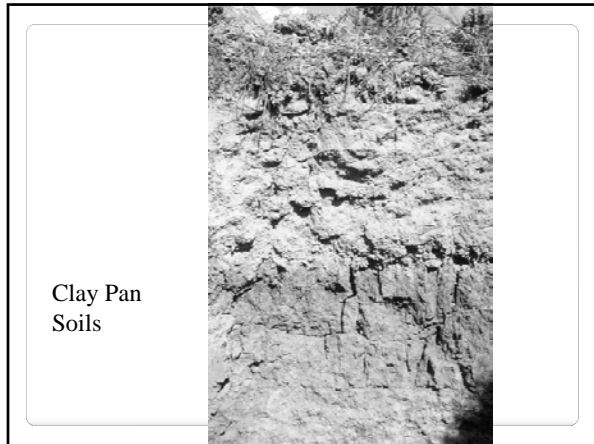
- Does water moves through all soils at the same rate?
- Soil horizons and restrictive features.
- Water tables.
- Interpretations: Hydraulic groups & Drainage classes.

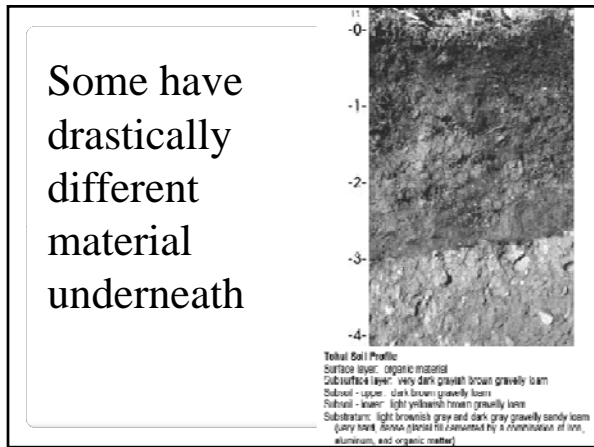
Soil Water Concepts

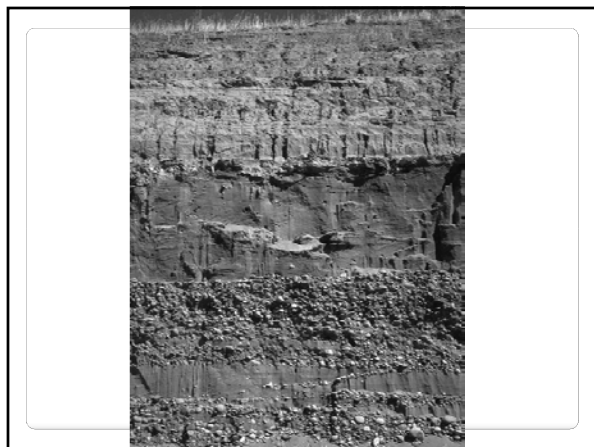


- clay pans
- abrupt texture change
- discontinuities
- bedrock

Soil Horizons and Restrictive Features







WATER TABLES

PERCHED (on something in the soil profile like)
fragipans
clay pans
abrupt texture change
discontinuities
bedrock

APPARENT (flood plains)
from the bottom up



Redoximorphic Features



The shapes and forms of bodies resulting from the **reduction** and **oxidation** of iron



Hydrologic Groups

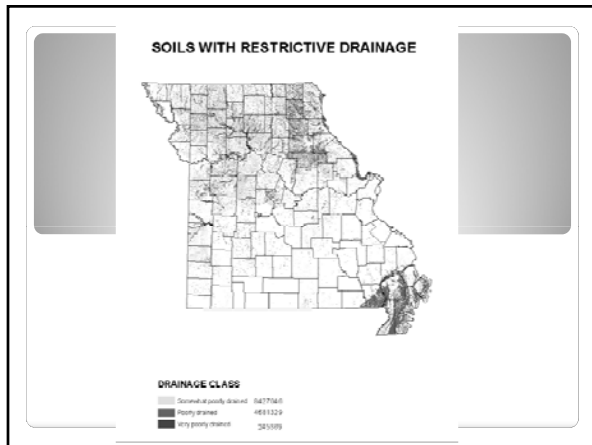
(a) Definition

Hydrologic group is a group of soils having similar runoff potential under similar storm and cover conditions.

(b) Classes

- The soils in the United States are placed into four groups, A, B, C, and D, and three dual classes, A/D, B/D, and C/D. In the definitions of the classes, infiltration rate is the rate at which water enters the soil at the surface and is controlled by the surface conditions. Transmission rate is the rate at which water moves in the soil and is controlled by soil properties. Definitions of the classes are as follows:
 - A.** (Low runoff potential). The soils have a high infiltration rate even when thoroughly wetted. They chiefly consist of deep, well drained to excessively drained sands or gravels. They have a high rate of water transmission.
 - B.** The soils have a moderate infiltration rate when thoroughly wetted. They chiefly are moderately deep to deep, moderately well drained to well drained soils that have moderately fine to moderately coarse textures. They have a moderate rate of water transmission.
 - C.** The soils have a slow infiltration rate when thoroughly wetted. They chiefly have a layer that impedes downward movement of water or have moderately fine to fine texture. They have a slow rate of water transmission.
 - D.** (High runoff potential). The soils have a very slow infiltration rate when thoroughly wetted. They chiefly consist of clay soils that have a high swelling potential, soils that have a permanent high water table, soils that have a claypan or clay layer at or near the surface, and shallow soils over nearly impervious material. They have a very slow rate of water transmission.

- Dual hydrologic groups, A/D, B/D, and C/D, are given for certain wet soils that can be adequately drained. The first letter applies to the drained condition, the second to the undrained. Only soils that are rated D in their natural condition are assigned to dual classes. Soils may be assigned to dual groups if drainage is feasible and practical.



Definition.—“Drainage class” identifies the natural drainage condition of the soil. It refers to the frequency and duration of wet periods.

DRAINAGE CLASS

- Excessively Drained >6'
- Somewhat Excessively Drained >6'
- Well Drained 3.5 to 6'
- Moderately Well Drained 2-3.5'
- Somewhat Poorly Drained 1-2'
- Poorly Drained 0-1'
- Very Poorly Drained +1'

7 Drainage Classes used in MO

• Saturated hydraulic conductivity or Ksat
 Pertains to the amount of water that would move downward through a unit area of saturated in-place soil in unit time under unit hydraulic gradient.

• Natural Drainage Class
 Refers to the frequency and duration of wet periods similar to those under which the soil developed.

Ksat vs. DRAINAGE

- Air space size
- Air space quantity
- Changes impact physical processes in the soil. Such as adhesion, cohesion, adsorption, and surface tension.

Soil Air Space & Water Movement

- Compaction breaks down the natural soil structure.
- Compaction changes the air space size and quantity.

Soil Compaction

◦ Saturated flow (also called gravitational flow) occurs only under saturated conditions when the force of gravity is greater than forces holding water in the soil.

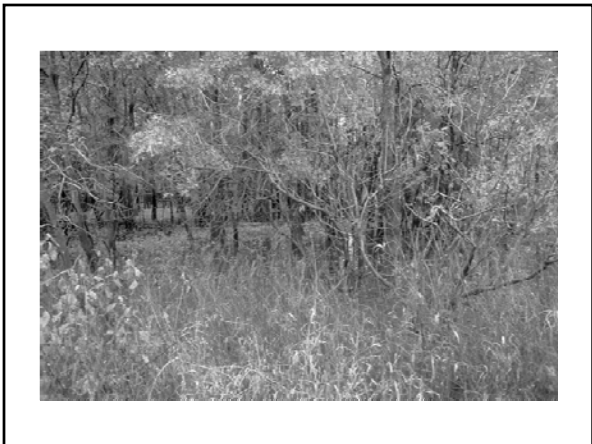
Capillary flow occurs in unsaturated soil (also called unsaturated flow). Unbroken films of water spread through connected capillary pores.

Flow terminology









DATA FORM
ROUTINE WETLAND DETERMINATION
(1987 C.D.R. Wetlands Information Manual)

Project/Site: [] Date: []
Applicant/Owner: [] County: []
Site number: [] State: []
Do Normal Circumstances exist on the site? Yes No Community ID: []
Is the site significantly disturbed (Appropriation)? Yes No Treatment ID: []
Is the area a special of Public Land? Yes No Plot ID: []
(If needed, explain on reverse)

VEGETATION

Domest Plant Species	Percent	Indicator	Domest Plant Species	Percent	Indicator
1	00	000	5	00	000
2	00	000	10	00	000
3	00	000	11	00	000
4	00	000	12	00	000
5	00	000	13	00	000
6	00	000	14	00	000
7	00	000	15	00	000
8	00	000	16	00	000

Percent of Domest Species that are OBL, FAC or PAC (excluding FAC): []
Remarks: []

HYDROLOGY

Recorded Data (Describe in Remarks):
Stream, Lake, or Tide Change
Aerial Photographs
Other: []
 No Recorded Data Available

Field Observations:
Depth of Standing Water: [](m)
Depth to First Water Table: [](m)
Depth to Saturated Soil: [](m)
Remarks: []

Wetland Hydrology Indicators:
Primary Indicators:
 Inundated
 Inundated as Upper 12 inches
 Water Marks
 Dark Layers
 Evidence of Soils
 Drainage Patterns in Wetlands
Secondary Indicators (2 or more required):
 Crustlike Soil Observed in Upper 12 Inches
 Wide Shallow Layers
 Local Soil Survey Data
 FAC Wetland Test
 Other (Explain in Remarks)

SOILS

Map Unit Name: [] Site ID: [] Plot ID: []
(Color and Photo) Drainage Class: []
Taxonomy (Soil type): [] Field Observations Confirm Mapped Type? Yes No

Profile Description:

Depth (cm)	Horizon	Mottled Color (Munsell Moist)	Moist Color (Munsell Moist)	Mott. Abundance / Contrast	Texture, Consistence, Structure, etc.
[]	[]	[]	[]	[]	[]
[]	[]	[]	[]	[]	[]
[]	[]	[]	[]	[]	[]
[]	[]	[]	[]	[]	[]

Hydro Soil Indicators:
 Stratified
 Blocky Structure
 Substratum
 Layered Structure
 Rooting Conditions
 Clayed in Low-Classes Colors
 Crustlike
 High Organic Content in Surface Layer in Sandy Soils
 Organic Structure in Sandy Soils
 Layered on Local Survey Data
 Based on National Hydro Soil Test
 Other (Explain in Remarks)

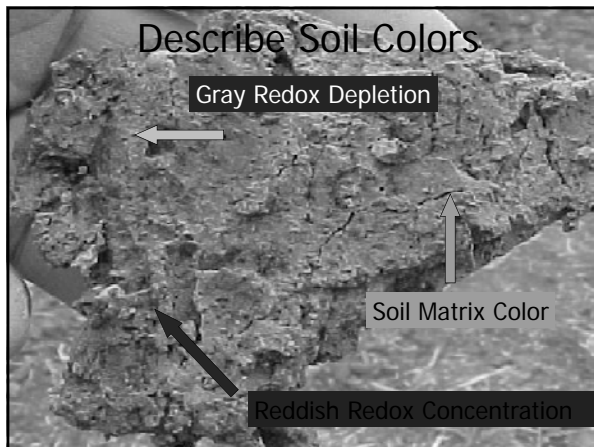
Remarks: []

Redoximorphic Features



The shapes and forms of bodies resulting from the **reduction** and **oxidation** of iron



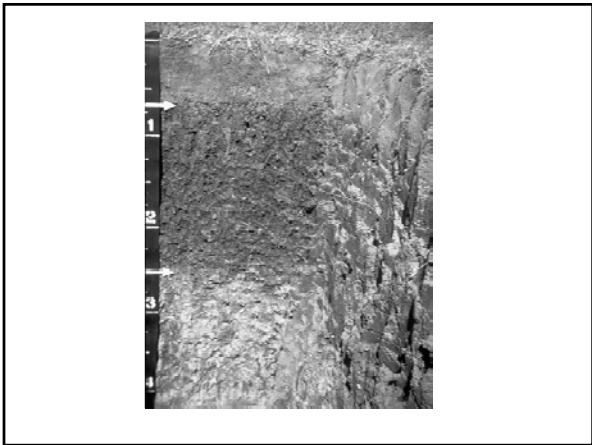


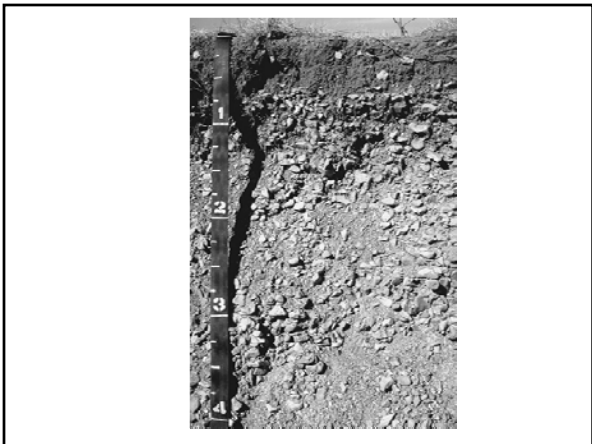


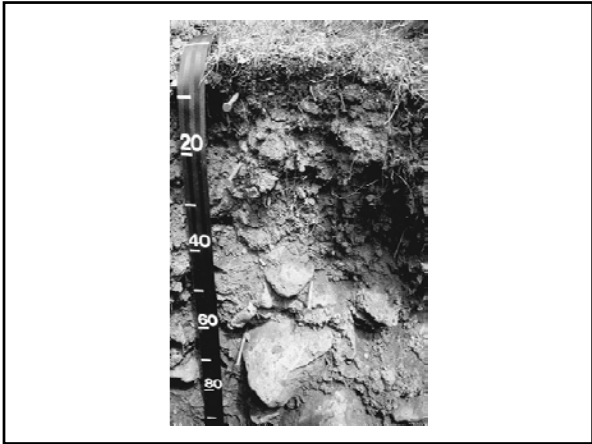












(1) Auger-hole method

The auger-hole method is the simplest and most accurate way to determine soil permeability (fig. 1). The measurements obtained using this method are a combination of vertical and lateral conductivity, however, under most conditions, the measurements represent the lateral value. The most limiting obstacle for using this method is the need for a water table within that part of the soil profile to be evaluated. This limitation requires more intensive planning. Tests must be made when a water table is available during the wet season. Obtaining accurate readings using this method requires a thorough knowledge of the procedure.

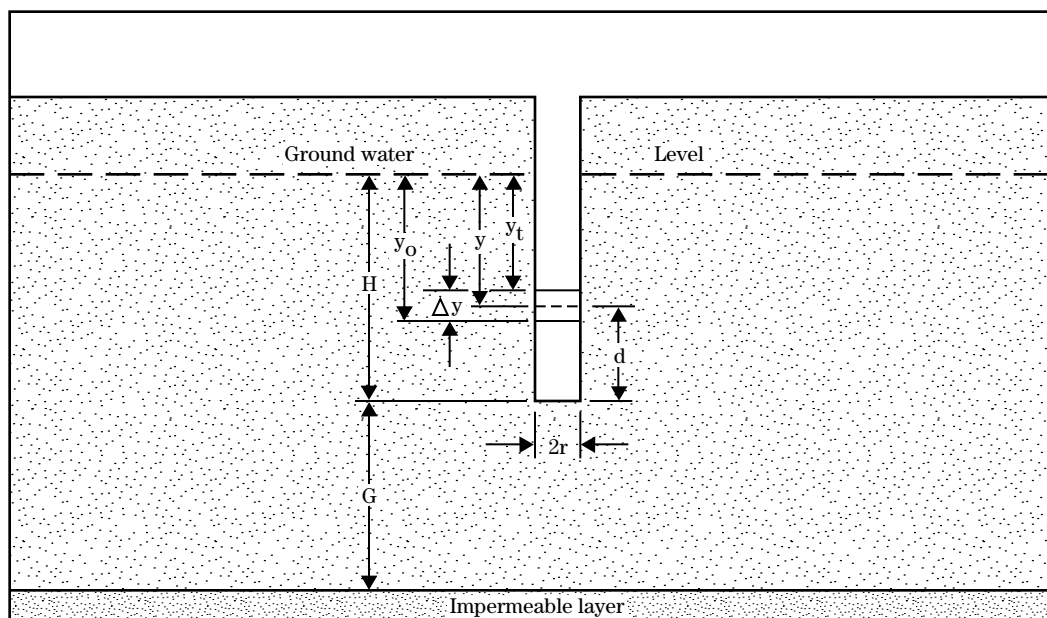
The principle of the auger-hole method is simple. A hole is bored to a certain distance below the water table. This should be to a depth about 1 foot below the average depth of drains. The depth of water in the hole should be about 5 to 10 times the diameter of the hole. The water level is lowered by pumping or bailing, and the rate at which the ground water flows back into the hole is measured. The hydraulic conductivity can then be computed by a formula that relates the geometry of the hole to the rate at which the water flows into it.

(i) Formulas for determination of hydraulic conductivity by auger-hole method—Determination of the hydraulic conductivity by the auger-hole method is affected by the location of the barrier or impermeable layer.

A barrier or impermeable layer is defined as a less permeable stratum, continuous over a major portion of the area and of such thickness as to provide a positive deterrent to the downward movement of ground water. The hydraulic conductivity of the barrier must be less than 10 percent of that of the overlying material if it is to be considered as a barrier. For the case where the impermeable layer coincides with the bottom of the hole, a formula for determining the hydraulic conductivity (K) has been developed by Van Bavel and Kirkham (1948).

$$K = \left(\frac{2220r}{SH} \right) \left(\frac{\Delta y}{\Delta t} \right) \quad [1]$$

Figure 1 Symbols for auger-hole method of measuring hydraulic conductivity



where:

- S = a function dependent on the geometry of the hole, the static depth of water, and the average depth of water during the test
 K = hydraulic conductivity (in/hr)
 H = depth of hole below the ground water table (in)
 r = radius of auger hole (in)
 y = distance between ground water level and the average level of water in the hole (in) for the time interval t (s)
 Δy = rise of water (in) in auger hole during Δt
 t = time interval (s)
 G = depth of the impermeable layer below the bottom of the hole (in). Impermeable layer is defined as a layer that has the permeability of no more than a tenth of the permeability of the layers above.
 d = average depth of water in auger hole during test (in)

A sample form for use in recording field observations and making the necessary computations is illustrated in figure 2. This includes a chart for determining the geometric function S for use in the formula for calculation of the hydraulic conductivity.

The more usual situation is where the bottom of the auger hole is some distance above the barrier. Formulas for computing the hydraulic conductivity in homogeneous soils by the auger-hole method have been developed for both cases (Ernst, 1950). These formulas (2 and 3) are converted to English units of measurement.

For the case where the impermeable layer is at the bottom of the auger-hole, $G = 0$:

$$K = \frac{15,000r^2}{(H + 10r) \left(2 - \frac{y}{H}\right) y} \frac{\Delta y}{\Delta t} \quad [2]$$

For the case where the impermeable layer is at a depth $\geq 0.5H$ below the bottom of the auger hole:

$$K = \frac{16,667r^2}{(H + 20r) \left(2 - \frac{y}{H}\right) y} \frac{\Delta y}{\Delta t} \quad [3]$$

The following conditions should be met to obtain acceptable accuracy from use of the auger-hole method:

- $2r > 2 \frac{1}{2}$ and $< 5 \frac{1}{2}$ inches
- $H > 10$ and < 80 inches
- $y > 0.2 H$
- $G > H$
- $y < 1/4 y_o$

Charts have been prepared for solution of equation 3 for auger-holes of $r = 1 \frac{1}{2}$ and 2 inches. For the case where the impermeable layer is at the bottom of the auger hole, the hydraulic conductivity may be determined from these charts by multiplying the value obtained by a conversion factor f as indicated on figure 3.

Figure 2a Auger-hole method of measuring hydraulic conductivity—sheet 1 of 2

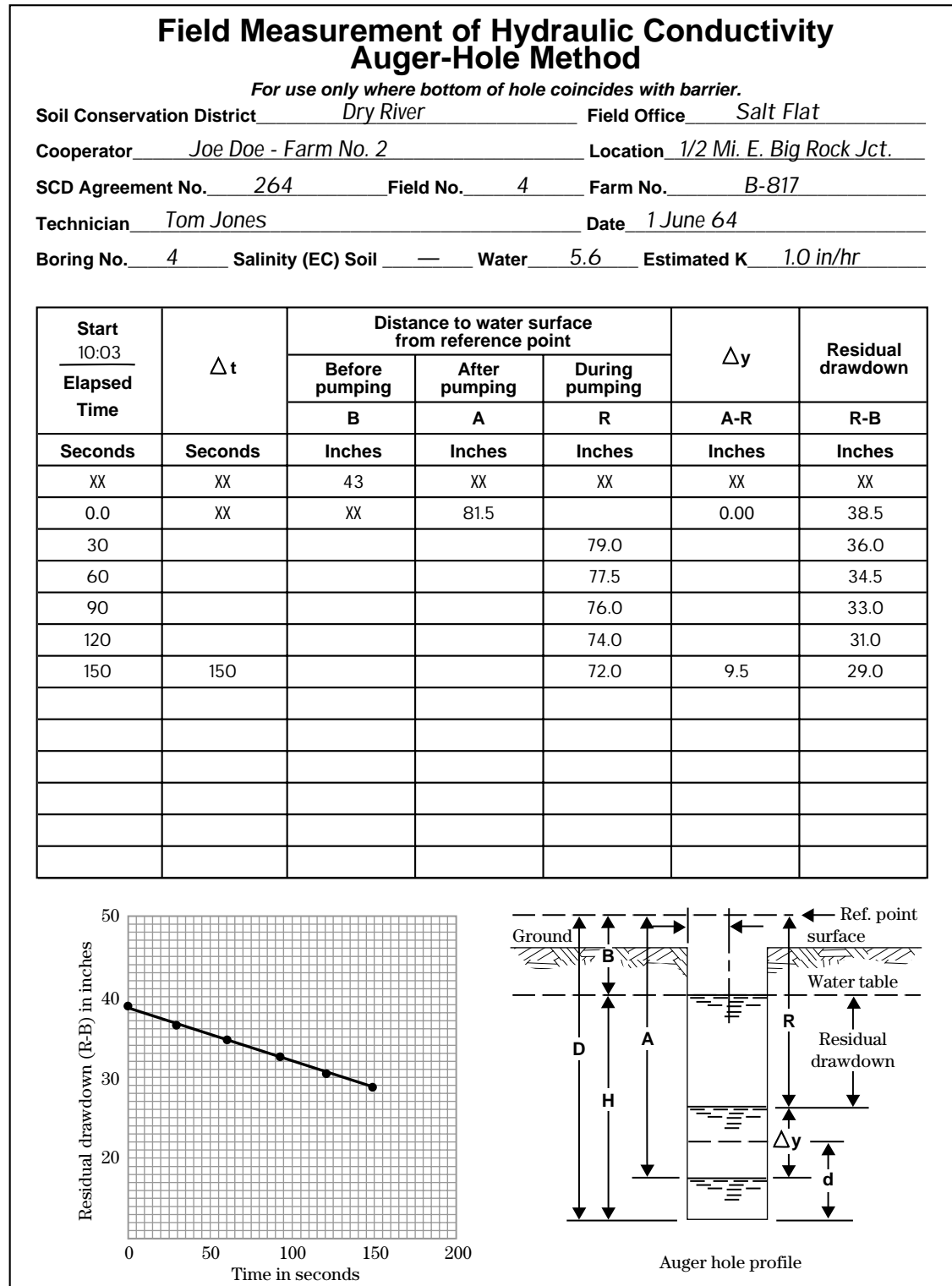


Figure 2a Auger-hole method of measuring hydraulic conductivity—sheet 2 of 2

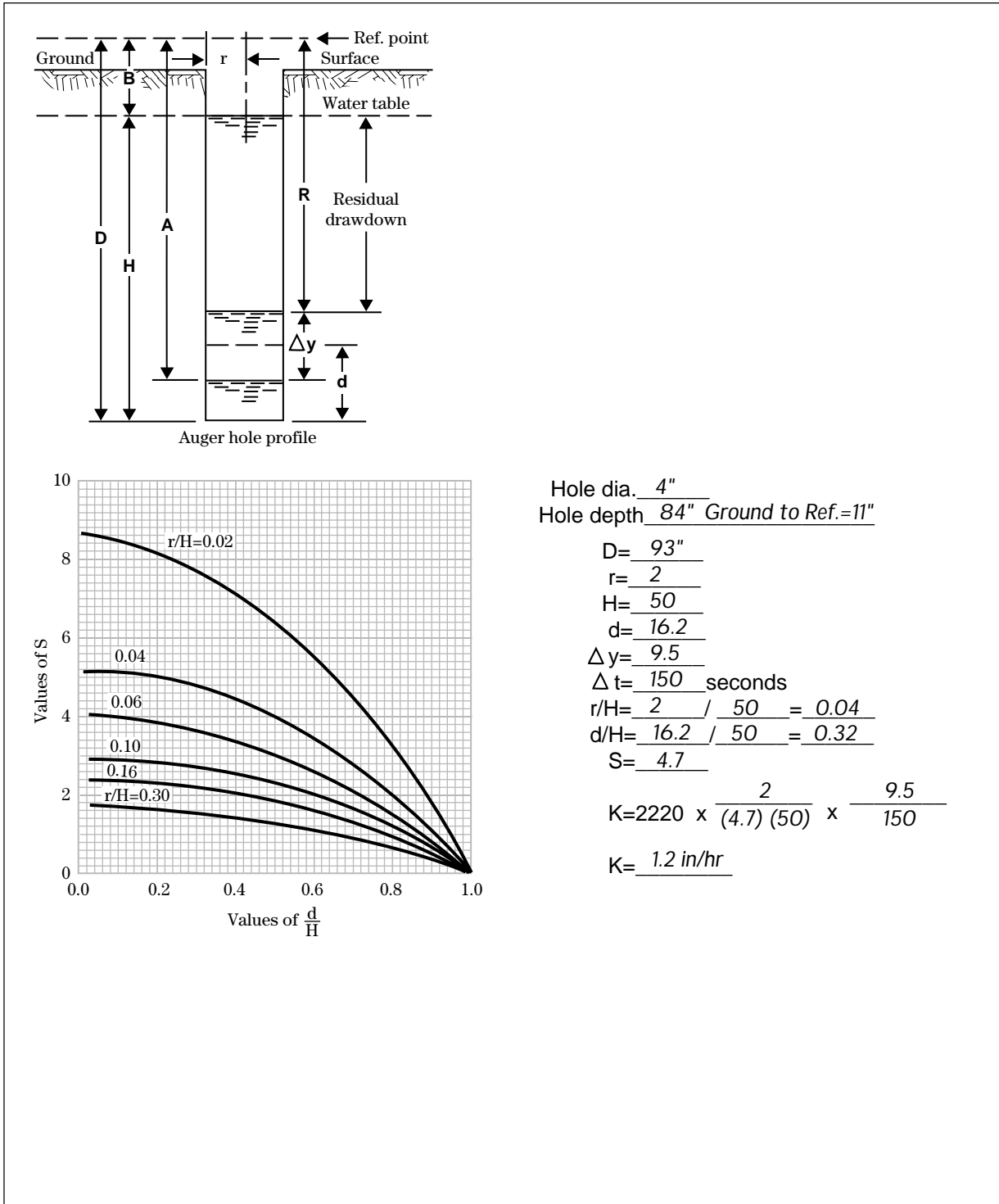


Figure 2b Auger-hole method of measuring hydraulic conductivity—sheet 2 of 2 (blank)

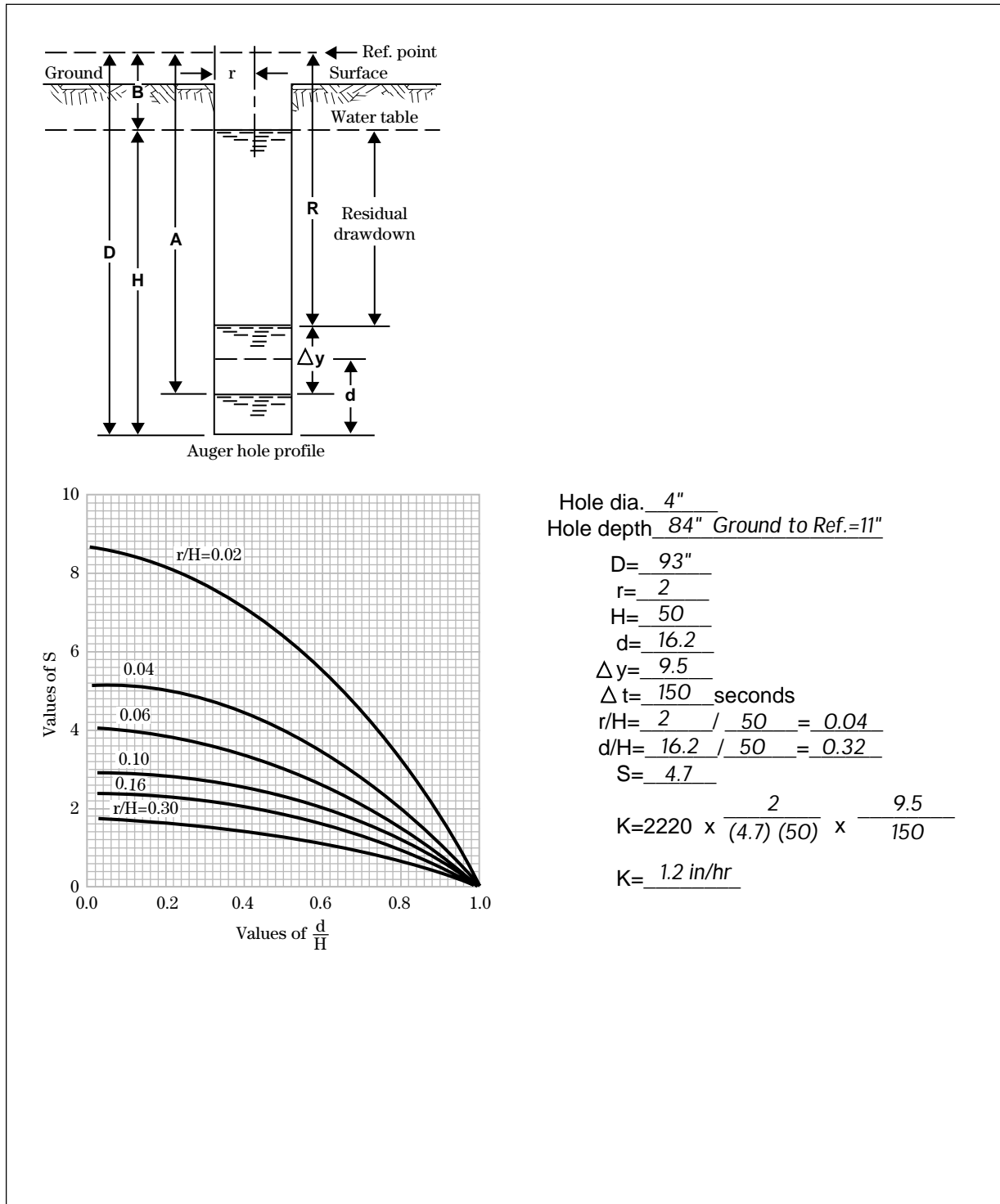


Figure 3 Hydraulic conductivity—auger-hole method using the Ernst Formula

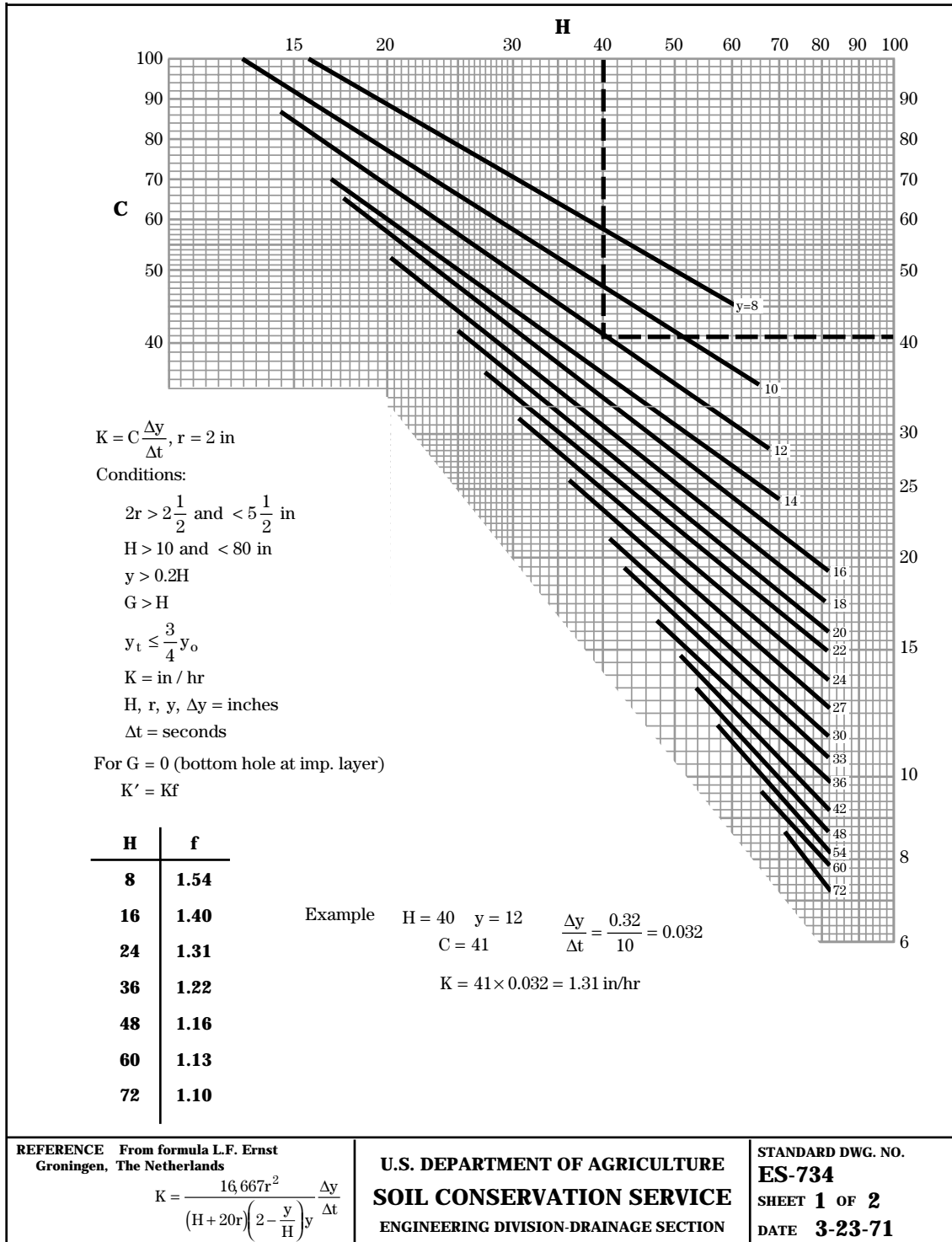
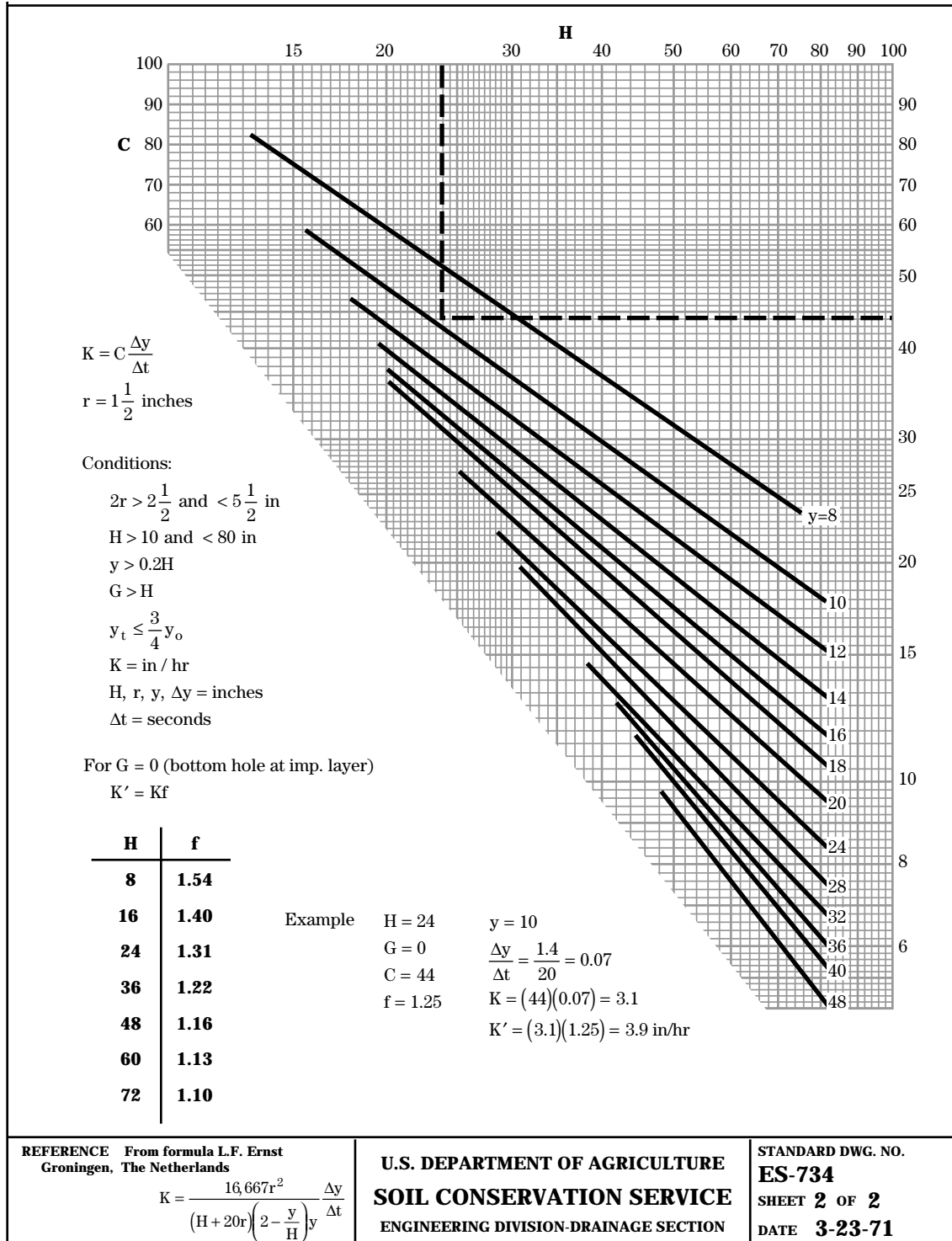


Figure 3 Hydraulic conductivity—auger-hole method by Ernst Formula—continued



(ii) Equipment for auger-hole method—The following equipment is required to test hydraulic conductivity:

- suitable auger
- pump or bail bucket to remove water from the hole
- watch with a second hand
- device for measuring the depth of water in the hole as it rises during recharge
- well screen may be necessary for use in unstable soils

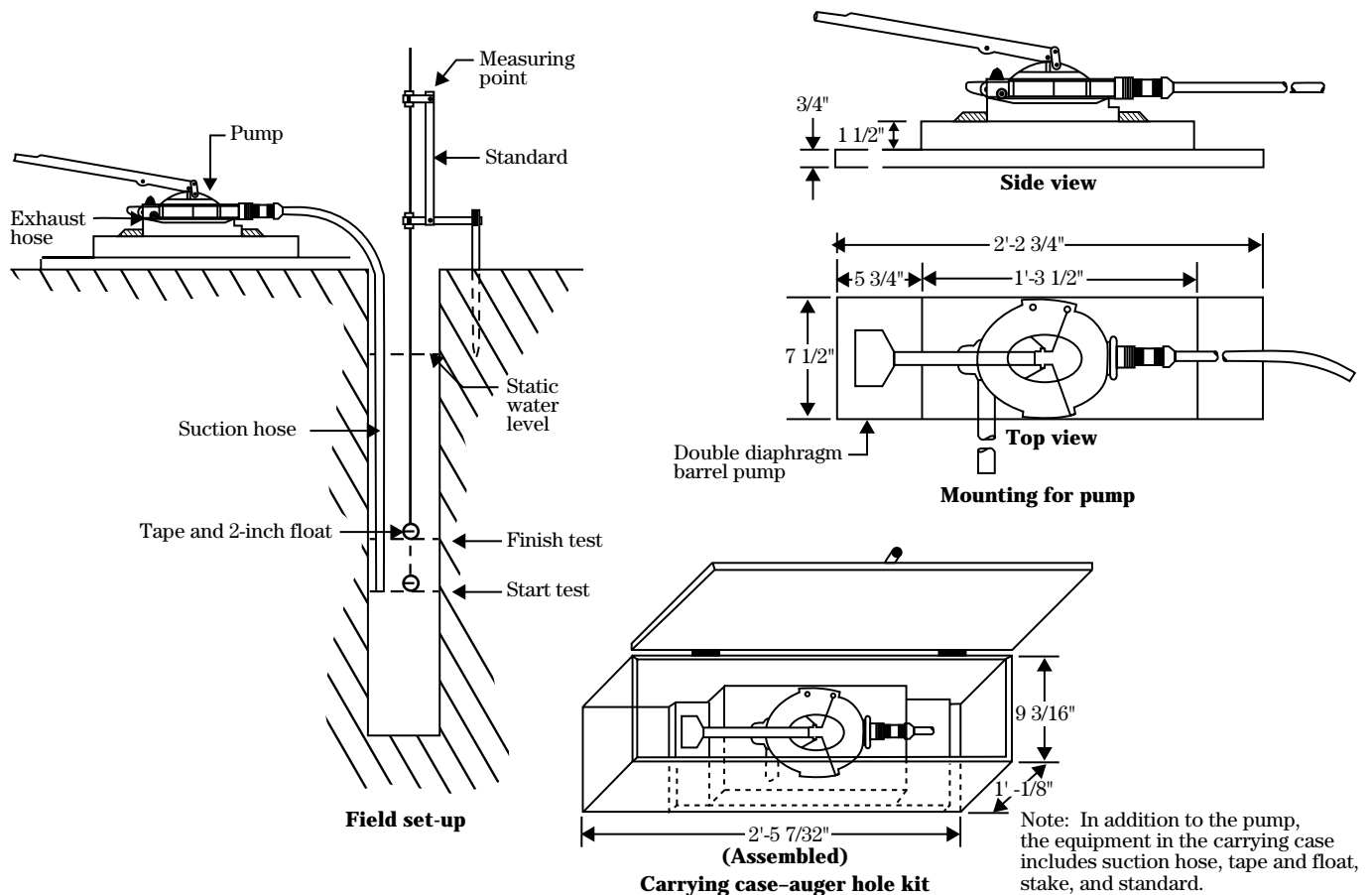
Many operators prefer a well made, light weight boat or stirrup pump that is easily disassembled for cleaning. A small, double diaphragm barrel pump has given good service. It can be mounted on a wooden frame for ease of handling and use.

For the depth measuring device, a light weight bamboo fishing rod marked in feet tenths and hundredths and that has a cork float works well. Other types of floats include a juice can with a standard soldered to one end to hold a light weight measuring rod.

A field kit for making the auger hole measurement of hydraulic conductivity is illustrated in figure 4. In addition to the items indicated in this figure, a watch and a soil auger are needed.

A perforated liner for the auger-hole is used in making the auger-hole measurement in fluid sands. This liner keeps the hole open and maintains the correct size. Several types of liners are used. Adequate slot openings or other perforations must be provided to allow free flow into the pipe.

Figure 4 Equipment for auger-hole method of measuring hydraulic conductivity



The openings in the screen should not restrict flow appreciably. The head loss through the screen should be negligible, and the velocity of flow through the openings should be small (0.3 foot per second or less) to prevent movement of fines into the hole. These criteria generally are met if the area of openings is 5 percent or more of the total screen area.

The Bureau of Reclamation uses 4-inch downspouting with 60 1/8- by 1-inch slots per foot of length. This works well in a variety of soils. A screen from the Netherlands is made from a punched brass sheet 2 millimeters thick with holes averaging about 0.5 millimeter in diameter. It is rolled into a tube 8 centimeters in diameter by 1 meter long. This screen works well because the sheet is rolled so that the direction in which the holes are punched is outward and the holes are variable in size. It has been used in many troublesome soils, and no clogging or failure to keep fines out of the hole has been reported.

Good judgment is needed in determining how far to drawdown the water level in the auger hole for the test. A minimum drawdown is necessary to physically satisfy theoretical criteria (refer to conditions given in fig. 3). Generally, a larger drawdown is made for slowly permeable soils than that for more permeable soils. A small drawdown for holes in sloughing soils may reduce the amount of sloughing. To prevent picking up sand in the pump, pumping should stop when the water level is within a few inches of the bottom of the hole.

Measurement of the rate of recovery of water in the auger hole should be completed before a fourth of the total amount of drawdown is recovered. Four or five readings should be taken at uniform short time intervals, and a plot of the readings made to determine a uniform rate of recovery to use in the formula. Plotting of time in seconds against the residual drawdown in inches indicates those readings at the beginning and end of the test that should be discarded and the proper values of t and y to use.

Drainage Design Reference

Area

1 acre = 43,560 square feet

Saturated Hydraulic Conductivity

1 micrometer per second = 1 $\mu\text{m}/\text{sec}$

1 $\mu\text{m}/\text{sec}$ = 0.2834 feet per day

1 $\mu\text{m}/\text{sec}$ = 0.1417 inch per hour

1 inch per hour = 7.0572 $\mu\text{m}/\text{sec}$

1 inch per hour = 2 feet per day

Pipe Flow

$$Q = V \times A \quad \text{and} \quad V = Q \div A$$

Where: Q = Flow discharge rate, cubic feet per second

V = Flow velocity, feet per second

A = Cross Sectional Area, square feet

Required Drainage Capacity

$$Q = 0.042 \times DC \times DA$$

Where Q = Flow discharge rate, cubic feet per second

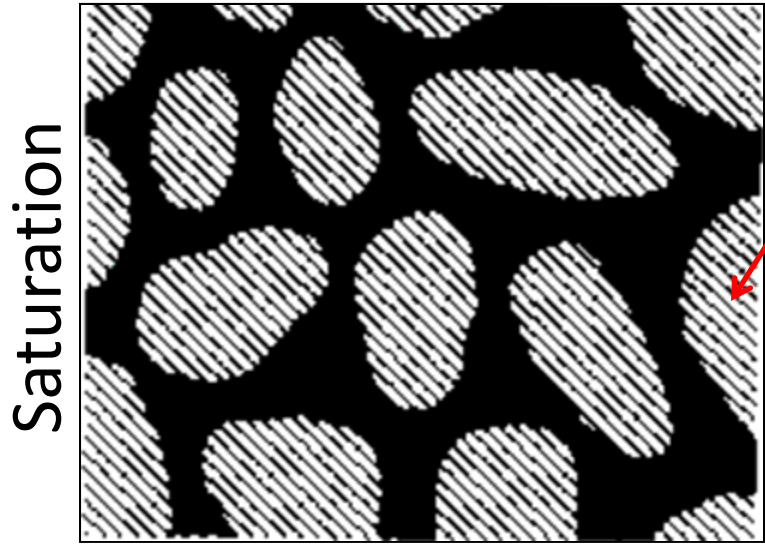
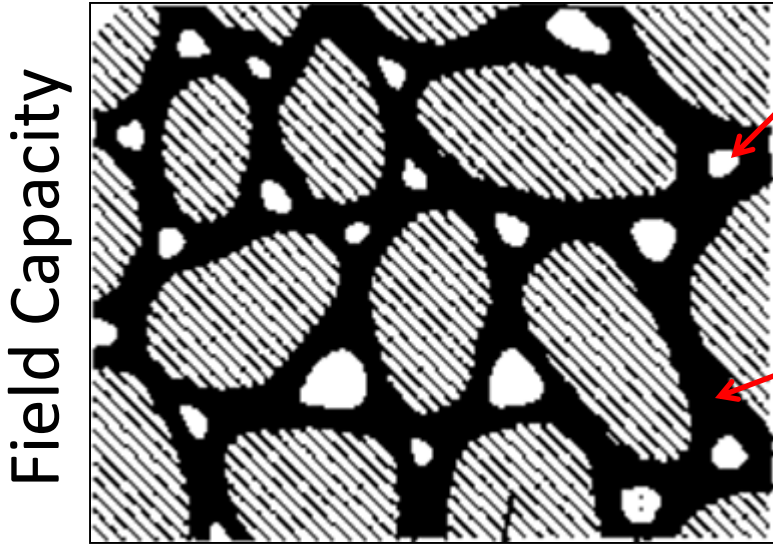
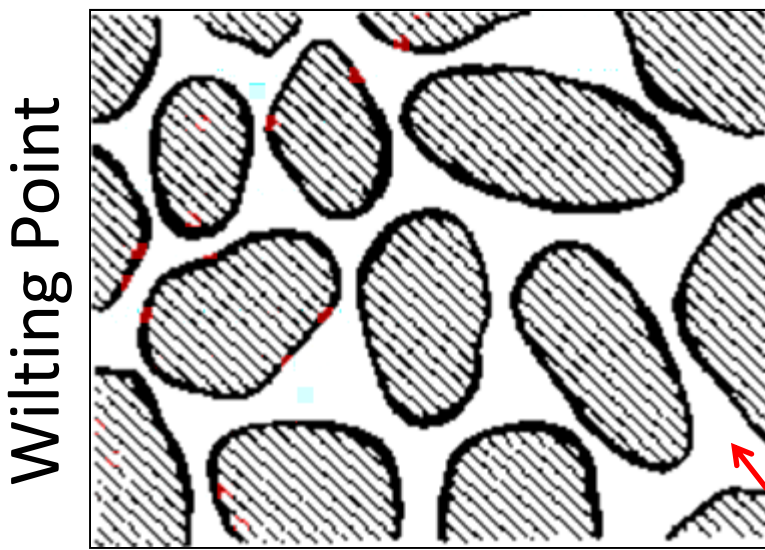
DC = Drainage Coefficient, inch/day

DA = Drained Area, acres

SATURATED HYDRAULIC CONDUCTIVITY - PERMEABILITY

K_{sat} Class		Permeability Class	
705.00	100.00	705.00	100.00
VERY HIGH		VERY RAPID	
		141.14	20.00
100.00	14.17		
HIGH		RAPID	
		42.34	6.00
10.00	1.417		
MODERATELY HIGH		MODERATELY RAPID	
		14.11	2.00
1.00	0.1417		
MODERATELY LOW		MODERATE	
		4.23	0.60
0.10	0.01417		
LOW		MODERATELY SLOW	
		1.41	0.20
0.01	0.001417		
VERY LOW		SLOW	
		0.42	0.06
0.00	0.00		
VERY SLOW		VERY SLOW	
		0.01	0.0015
IMPERMEABLE		IMPERMEABLE	
		0.00	0.00
µm/sec	in/hr	µm/sec	in/hr
<i>µm/sec x 0.1417 = in/hr</i>		<i>in/hr x 7.0572 = µm/sec</i>	
<i>µm/sec x 0.2834 = ft/day</i>		<i>in/hr x 2 = ft/day</i>	

Soil Water



Air

Water

Soil

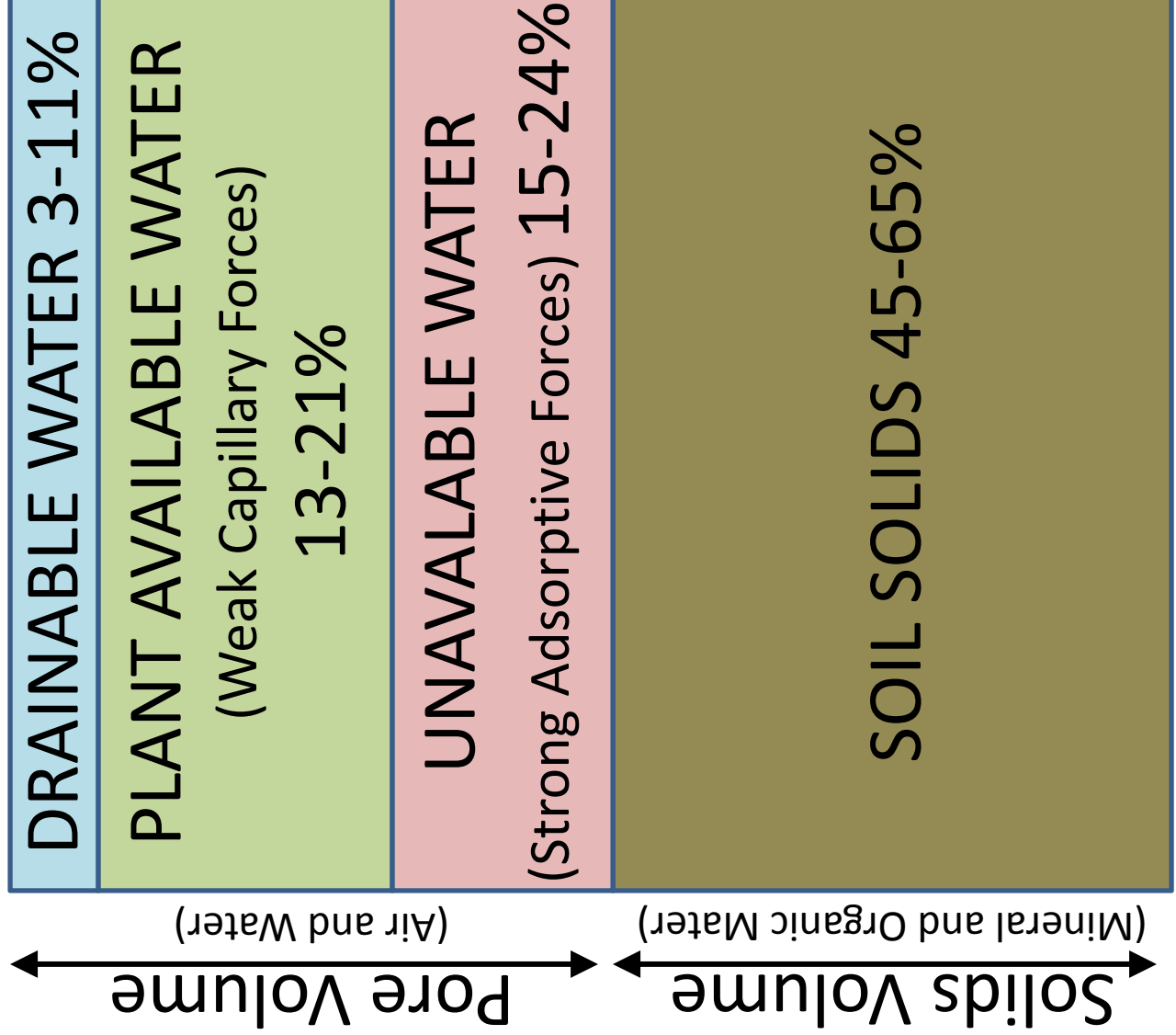


Available
Water



Drainable
Water

- ← Saturation
- ← Field Capacity
- ← Wilting Point
- ← Completely Dry

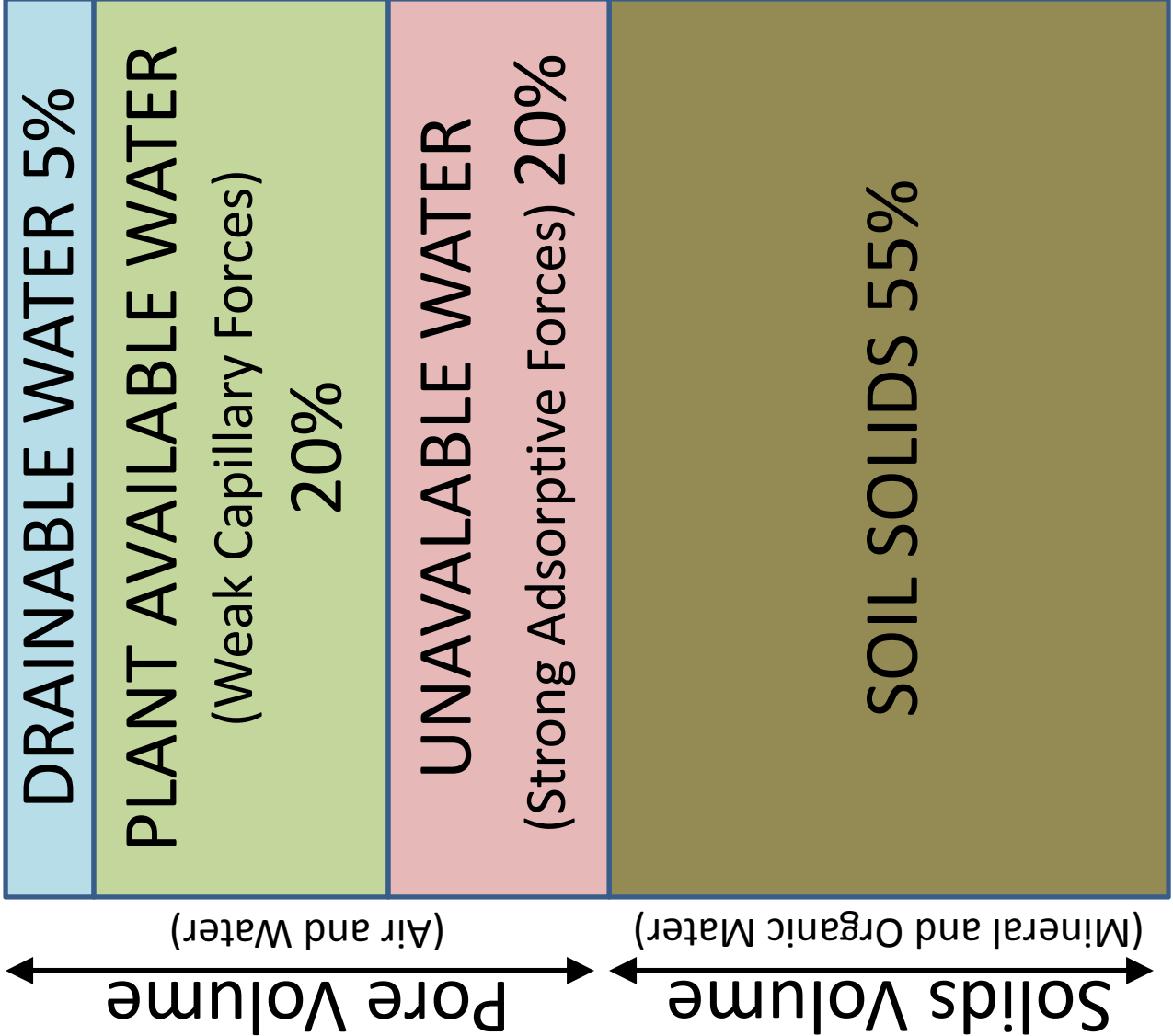


Typical Soil Water Relationships

Soil Texture	Wilting Point (% by vol.)	Available Water (% by vol.)	Drainable Water (% by vol.)
clays, clay loams, silty clays	15-24	15-26	3-11
well structured loams	8-17	12-22	10-15
sandy	3-10	7-20	18-35

Source: University of Minnesota BU-07644-S, Soil Water Concepts, Gary Sands

- ← Saturation
- ← Field Capacity
- ← Wilting Point
- ← Completely Dry



Soil Water Relationships
(Example Silty Clay)

How much water do I need to remove?

Given a soil (silty clay) with a drainable porosity of 5% with the goal of draining the top 12 inch layer in 48 hours.

Volume of drainable water

$$= 5\% \times 12 \text{ inch depth}$$

$$= 0.6 \text{ inches}$$

Rate of removal

$$= 0.6 \text{ inch} \div 2 \text{ day}$$

$$= \underline{0.3 \text{ inch/day}}$$

How much water do I need to remove?

Given a soil (loam) with a drainable porosity of 12% with the goal of draining the top 12 inch layer in 48 hours.

Volume of drainable water

$$= 12\% \times 12 \text{ inch depth}$$

$$= 1.4 \text{ inches}$$

Rate of removal

$$= 1.4 \text{ inch} \div 2 \text{ day}$$

$$= \underline{0.7 \text{ inch/day}}$$